are seen on the surface of the delta in the sand dunes which have there developed.

Along the southeastern border of the delta the land drops off to a clayey terrace lying between 250 and 280 feet in elevation. It is well exhibited at Reynolds Corners [*see* pl. 13]. The slope from the 350 foot delta plain to this lower terrace from Fort Edward southward coincides closely with the boundary between the Calciferous-Trenton limestones and the Hudson river shales which lie on the east of them. But the immediate origin of the slope appears to be due to erosion taking place subsequent to the formation of the 350 foot delta. This lower terrace corresponds in position with a tilted water plane of a glacial lake whose outlet on the south, as is shown on plate 28, was in the

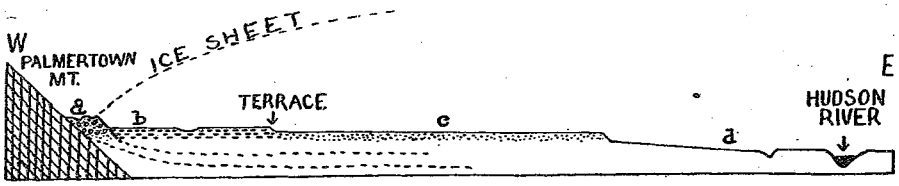

Fig. 20 Profile of the terraces and delta levels from the base of Palmertown mountain to the Hudson river. *a*, The glacial terrace; *b*, the broader terrace at 400 feet; *c*, the delta of the Hudson; *d*, the clay terrace, a part of the channel of the stream which flowed through the Coveville outlet

old channel back of Schuylerville which falls into the Hudson gorge at Coveville.

Below the level of this terrace is the old channel continuing the Hudson gorge by way of the Wood creek valley to Lake Champlain, the evident path of a river which as I shall hope to show later in this report drained a glacial lake in the Champlain valley into the Hudson gorge. East of this channel is a branch at a somewhat higher level perhaps earlier occupied by the same stream before that nearer Fort Edward was so deeply excavated.

The above diagram, figure 20, is intended to show by an east and west profile the successive terrace and delta levels of the Fort Edward district, down to the existing channel of the Hudson below the site of the old fort.

The effect of ice barriers and glacial lakes about the southeastern base of the Adirondacks is so well exhibited in the case of the Hudson river that the following digression is introduced