

presented itself, it might be expected that the ice would deploy into it. Second, these local striae are due to a branch of the main Champlain lobe which passed southward through the mountain valleys on the west and rejoined the main lobe at Port Henry. Third, the striae are due to a local glacier or to small glaciers coming at least to within 60 feet of the present lake level from snow fields in the high area on the west, whose culminating point, Mt Marcy, is 24 miles distant in a west by north direction.

The position and nature of the till deposits of morainal character shown on the map are explicable by the second and third hypotheses. The fact that the eastwest striae extend quite down to present lake level favors the first hypothesis on account of the difficult assumption that a local glacier would push its front to so low a level after the main ice had retreated from this latitude. On the other hand the striae near the present lake shore come to the shore so abruptly as to give no support to the first hypothesis. I was not able to make out from the details of the few striae observed whether the ice movement indicated by them was toward or away from the present lake. In the case of the ledge near the red schoolhouse southwest of Port Henry there are crescentic flakes in the gneissoid rock with the horns pointing to the northeast. I have seen similarly fractured pieces removed in which the horns of the crescent pointed in the direction in which the ice was moving. If that be true here it is a point in favor of local glaciation.

In the Mill brook locality, the glaciated ledge of the abnormal set was overlain by a gravelly boulder clay containing fragments of Potsdam sandstone, gneisses and limestones.

Bulwagga mountain comes to the lake front with a bold escarpment and, blocklike, rises between the depressed region of Port Henry and the similar area of Crown Point. In this latter side valley glacial striae, n. 42° e., were observed just south of the village on the north slope of Sugar hill. On the road from Crown Point Centre to Coot hill, glacial striae may be seen running n. 48° e.

Yet farther south of the west side of the lake in the vicinity of Ticonderoga, the striae range from n. 13° e. to n. 33° e., dis-