total amount of wear since Utica time has been exceedingly small, and the surface must have been for much of the time near base level. Moreover, such blocks were the most downfaulted of all and must have formed depressed basins with every period of renewed faulting, as such receiving deposit from the surrounding higher blocks as these were worn down, thus protecting their own surfaces from wear for long intervals.

This erosion interval was so protracted, extending through the greater part of Mesozoic time, that the whole region was finally planed down to a surface of little relief, broad shallow valleys and low divides, with occasional low, rounded hills or clump of hills where extraresistant rocks occurred, or where favorable location prevented maximum wear. The fault cliffs, or scarps, were entirely wiped out as topographic features, the raised side being worn down to the same level as the dropped.

In the southern and western Adirondack region this old surface is still recognizable, as the upland surface into which the present valleys are cut, the old residual hills rising above it now as they then did. The present plateau upland of southeastern New York would seem to represent a continuation southward of this same old surface; and, if this be the case, the result in the Adirondacks was simply a local development of conditions which prevailed widely in the eastern United States at this time.

Cenozoic history

This long period of quiet wear was terminated by another uplift, which would seem to have occurred at the close of Mesozoic, or the beginning of Cenozoic time. This uplift inaugurated another erosion cycle with a much lower base level. In the Adirondack region this uplift was, at least for the northern part of the district, of a dome-shaped character, the major axis of the dome being along a nearly north-south line closely coinciding with the line between Clinton and Franklin counties, thence turning southwestward; the minor axis running through the extreme south of Franklin county and thence eastward through Essex. In the eastern Adirondacks this uplift was complicated by further shifting along the fault planes, bringing fault cliffs, or scarps, again